

A climate change scenario and soil ammonium “fixation” during the seasonal rice (*Oryza sativa*) growth in Portugal under intermittent flooding

A mudança climática na “fixação” do amónio no solo durante o ciclo cultural de arroz (*Oryza sativa*) sob alagamento intermitente, em Portugal

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ABSTRACT

The “newly fixed” NH_4^+ in clay minerals should be considered for an efficient management of plant nutrition. In a clay-loam soil cultivated in 2012 with rice under intermittent flooding and conventional agronomic practices, the dynamics of pH, N-inorganic and non-exchangeable NH_4^+ was evaluated *i*) under field conditions, air temperature and atmospheric $[\text{CO}_2]$ ($375 \mu\text{mol mol}^{-1}$ air) in Salvaterra de Magos (Portugal), and *ii*) under elevated temperature and temperature+ $[\text{CO}_2]$ in controlled microclimate. For that, open-top chambers were used to simulate the conditions of an increased mean air temperature ($2\text{-}3^\circ\text{C}$), alone or combined with increased $[\text{CO}_2]$ ($550 \mu\text{mol mol}^{-1}$ air). Non-exchangeable NH_4^+ was significantly higher under open-field conditions compared with the temperature elevation, with or without elevated atmospheric $[\text{CO}_2]$. Temperature elevation reduced the “fixation” rate of the cation, while the CO_2 concentration rise did not affect particularly the non-exchangeable form. Further studies are required for consolidation of these findings together with microbial communities and dynamics of C and N in soil.

Keywords: Atmospheric CO_2 , clay loam soil, controlled environment, open-field, temperature.

RESUMO

O NH_4^+ recém-“fixado” nos minerais argilosos deve ser considerado numa gestão eficiente da nutrição vegetal. Num solo argilo-limoso, localizado em Salvaterra de Magos (Portugal), cultivado em 2012 com arroz, sob alagamento intermitente e práticas agronómicas convencionais, avaliou-se a dinâmica do pH, N-inorgânico e NH_4^+ não-trocável *i*) em condições de campo, de temperatura do ar e $[\text{CO}_2]$ atmosférico ($375 \mu\text{mol mol}^{-1}$), -, e de *ii*) aumento da temperatura e temperatura+ $[\text{CO}_2]$ em microclima controlado. Foram utilizadas câmaras de topo aberto para simulação do aumento da temperatura média do ar ($2\text{-}3^\circ\text{C}$), apenas ou combinado com aumento da $[\text{CO}_2]$ ($550 \mu\text{mol mol}^{-1}$). Os resultados mostraram que o NH_4^+ “fixado” foi superior em condições de campo, em comparação com a elevação da temperatura, com ou sem elevação da $[\text{CO}_2]$. A elevação da temperatura reduziu a taxa de “fixação” do catião, enquanto o aumento da $[\text{CO}_2]$ não afectou o NH_4^+ “fixado”. Mais estudos são necessários para a consolidação destes resultados, bem como da dinâmica do C, N e população microbiana no solo.

Palavras-chave: Ambiente controlado, ar livre, concentração de CO_2 , solo argilo-limoso, temperatura.

Introduction

Rice is a staple food for nearly half of the world's population, most living in developing countries. Portugal is the first rice consumer *per capita* in Europe and the fourth producer (6 t ha⁻¹ in 28,000 ha), contributing to the 5.3% of the total European production. Rice cultivation in Portugal is mostly located in the central and southern regions (Mondego, Tagus and Sado valleys) as a monoculture under flooding or intermittent flooding regime. The height of floodwater varies considerably during the season depending on plant size, climatic conditions, soil type and agronomic practices.

Simultaneous supply of ammonium (NH₄⁺) and nitrate (NO₃⁻) can improve rice growth (Shi *et al.*, 2010), but NH₄⁺ is the preferred form of nitrogen (N) for an efficient rice absorption and microbial assimilation under flooded conditions. Ammonium transformation in the paddy rice field is an important process with respect to both agricultural production and environmental protection. Flooding enhances ammonification by facultative anaerobic microorganisms, providing NH₄⁺ to the soil solution. This cation is then adsorbed or "fixed" (non-exchangeable form) in the clay minerals, especially the 2:1 expansive minerals (e.g. illite, vermiculite, smectite). This non-exchangeable cation in soils rich in clay minerals contributes markedly to rice nutrition. The effects of fertilization on NH₄⁺ sequestration in soils have been reported mostly for upland unflooded soils (e.g. Carranca, 1996). Few studies have been made on paddy soils, and the underlying mechanisms for NH₄⁺ exchange are far less understood with contradictory reports (Scherer and Zhang, 1999; Liu *et al.*, 2008).

Ammonium is captured by the contracted interlayer's of clay minerals after addition of N fertilizers ("newly fixed" NH₄⁺) and becomes temporarily unavailable for crop nutrition. In flooded rice soils, the anoxic conditions following soil submergence and the accumulation of NH₄⁺ are favorable for the temporary "fixation" of this N form in soils with high amounts of expandable 2:1 clay minerals. The "newly fixed" NH₄⁺ is thus protected from N losses *via* nitrification, denitrification and volatilization processes which may occur during the drying and rewetting of paddy rice soils under the intermittent water regime. In waterlogged soils, NH₄⁺ "fixation" is stimulated by flooding because soil organic carbon enhances the decline in redox potential (Eh) (Schneiders and Scherer, 1998) which induces a reduction of octahedral iron (Fe) in 2:1 clay minerals, and therefore an increase in the nega-

tive charge with a simultaneous higher Coulombic attraction, i.e. higher affinity between the interlayer cations and the silicate layers (Scherer and Zhang, 1999, 2002; Liu *et al.*, 2008). Therefore Eh may have an important impact on the fate of NH₄⁺ in paddy soils. Further, coating of the surface of clay minerals by Fe oxides also has an impact on the diffusion of NH₄⁺ into or out of the interlayers of clay minerals (Scherer and Zhang, 1999). Therefore, the reversible oxidation and reduction of Fe oxides in paddy rice soils is a mechanism of special importance in sorption/"fixation" and desorption/diffusion of NH₄⁺ and its availability to rice.

The rate of "fixation" is rapid, with the proportion of "fixed" NH₄⁺-N:total N ≥ 11% in cultivated soils (Carranca, 1996; Qiu *et al.*, 2012). The "recently fixed" NH₄⁺ can be released into the soil solution when the clay interlayer's expand and the NH₄⁺ in the soil solution falls below a certain level, becoming available to plants and soil microorganisms. The release of NH₄⁺ is highest in the rhizosphere of rice plants, by the presence of oxygen (O₂) and/or NO₃⁻ and decreases with distance from the roots, correlating well with the Eh (Zhang and Scherer, 2002). The "fixation" and subsequent release of "recently fixed" NH₄⁺ from fertilizer N is of importance in its uptake by plants and differs remarkably with soil texture, clay mineral composition, pH, temperature, fertilization practice, and cropping system (Carranca, 1996; Lu *et al.*, 2010). Results on the availability of non-exchangeable NH₄⁺ for crops are contradictory and are still open to discussion.

The climate change has increased the atmospheric carbon dioxide concentration ([CO₂]) from 280 to 380 μmol mol⁻¹ of air in the past two centuries, and is expected to reach 700-1000 μmol mol⁻¹ of air until 2100 if no mitigation measures to reduce CO₂ emissions are accounted for (IPCC, 2007). In addition, the rise in the global average air temperature will be between 1.8 to 4.0 °C by the end of the present century relative to the mean value for 1980-1999. Scenarios of climate change by temperature and CO₂ elevation can be particularly relevant in Mediterranean regions. The continuing increase in atmospheric [CO₂] and projections of possible future increases in global air temperature caused by the [CO₂] elevation have stimulated scientific interest in the effects of these climate variables on important food crops, but literature on soil N, C and microbial dynamics is missing. As to the temperature effect *per se* literature is contradictory sometimes reporting that "fixation" of NH₄⁺ in soils increases with the temperature rise by the dehydration and contraction of the crystal lattice. Juang (1990) observed that treatments with

temperatures at 25 °C, 40 °C and 70 °C had little influence on the non-exchangeable NH_4^+ in both the soil and clay, while at 110 °C, the “fixed” cation in the clay increased significantly, by 13% in a Latosol and 99% in a recent sandstone-shale alluvial soil. This indicates that a temperature of 110 °C may remove more water molecules from the interlayer of the clay minerals, thus increasing the amount of “fixed” NH_4^+ in the clay fraction. By contrast, other authors related that “fixation” is stronger below 30 °C, and even more by freezing. Jia *et al.* (2000) showed that the “fixation” rate of the fertilizer- NH_4^+ was 18%-23% at normal temperature (25 °C) and 38%-45% at low temperature (-5 °C to 0 °C) under incubation condition.

In the present study we hypothesized that the climate change (temperature and CO_2 elevation) may affect NH_4^+ “fixation” due to changes in soil physico-chemical and mineralogical properties. In a clay loam soil, at Salvaterra de Magos (central Portugal), cultivated in 2012 with rice (*Oryza sativa* L. cv. Ariete) under intermittent flooding and conventional agronomic practices, the dynamics of pH, inorganic N and non-exchangeable NH_4^+ were evaluated for *i*) the natural ambient conditions (open-field), and *ii*) the temperature and CO_2 +temperature increase in controlled environments in field.

Material and Methods

Site description and layout of field experiment

A field experiment with a *japonica* rice variety (*Oryza sativa* L. cv. Ariete) was conducted in 2012 at Salvaterra de Magos (Tagus valley, central Portugal; latitude: 39°2'20.15"N, longitude: 8°44'25.7"W, elevation: 18 m above sea level). Rice was sown on 23 May 2012 at a rate of 200 kg dry seeds ha^{-1} in a clay loam soil with intermittent flooding regime. The study area is the main region for rice production in Portugal. The experimental design consisted of three treatments arranged in a randomized complete block design (RCBD) with three replicates, in a total of nine blocks (Figure 1a). Each block was 4.0 m x 4.0 m, separated 4 m from each other. Treatments were as follows: elevated CO_2 +temperature, elevated temperature, and the unchambered control plots with the natural ambient conditions (CO_2 concentration of $375\pm 46 \mu\text{mol mol}^{-1}$ air, and ambient air temperature and rainfall). To change the climate variables, six large open-top chambers (OTCs) (OTC=4 m wide x 3 m height x 2 m open-top diameter, 30° tilt) covered with a polyethylene film (1 mm thickness and 75% light transmittance and provided by

EstufasMinho, S.A., Fão, Portugal) except for the open-top (Figure 1b) were placed on a previous prepared (chisel and laser) lowland for paddy conditions: three OTCs for elevated CO_2 +temperature and three for the temperature rise by the chamber effect. Details on the construction and operation of OTCs have been provided by Pereira *et al.* (2013). In brief, in the three OTCs for CO_2 enrichment, a system using pure industrial CO_2 injection was installed to fumigate CO_2 during the day-night time (24 h per day). It operated from May to October 2012 in order to have an average atmospheric CO_2 concentration of $547\pm 73 \mu\text{mol mol}^{-1}$ air, which represented the expected range by the middle of the 21st century.

The Anthropic soil (World Reference Base for Soil Resources – FAO/ISRIC/IUSS, 2006) was representative for rice production in Portugal. It had a clay loam texture (174, 276 and 551 g kg^{-1} of sand, silt and clay, respectively) in the 0-60 cm layer. The dominant clay minerals in the soil were determined by the classical X-ray diffraction methods (Philips *et al.*, 1971) and followed the order: illite (200-500 g kg^{-1})>smectite (20-200 g kg^{-1})>vermiculite and interstratified with illite, vermiculite and smectite (<20 g kg^{-1}). In the surface 0-20 cm layer, the bulk density was 1.1 g cm^{-3} , $\text{pH}_{(\text{H}_2\text{O})}$ was 5.9, organic C was 24 g kg^{-1} , total N was 2.4 g kg^{-1} , and cation exchange capacity was $22.7 \text{ cmol}_{(+)} \text{ kg}^{-1}$.

During the growth season, a mineral N fertilizer was split and applied twice. All the experimental plots received the same doses of fertilizers. A NP mineral fertilizer (20-20-0) was mechanically broadcast into the soil in 14 May as a basal dressing, at a rate of 60 $\text{kg NH}_4\text{-N ha}^{-1}$ preceding the crop seeding, and a topdressing was carried out in 10 July, at tillering, in the form of sulfamid (40% N) and at rate of 60 kg N ha^{-1} . The basal fertilizer was incorporated into the amended soil with partially burnt crop residues (straw+root) at a depth of 20 cm, and the topdressing was applied by hand on the floodwater. No potassium was added to the soil as the soil was rich in this nutrient, whereas 60 kg of phosphorus per hectare was incorporated into the soil as part of the basal dressing.

After rice seeding, the water regime was intermittent, i.e. flooding - midseason drainage (for plant rooting, one week after rice germination, and twice for weed control during a few days at tillering) - reflooding - drainage (three weeks before crop harvest) (Figure 2). Floodwater height varied from 5 cm to 20 cm during the crop season in the open-field, as well as inside the OTCs by passing through some holes at the bottom of the polyethylene film involving the OTC. Irrigation water had a pH 8.0,

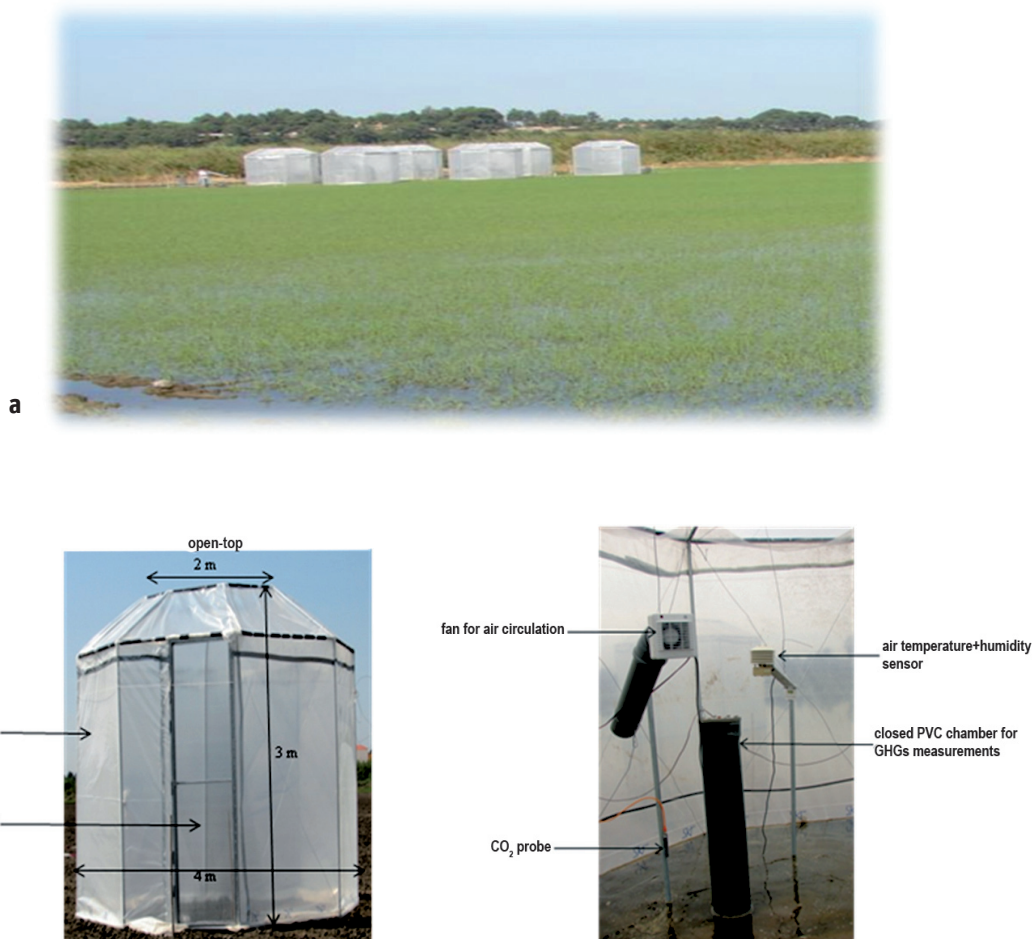


Figure 1 – A partial aspect of the experimental site (a) installed in a rice field at Salvaterra de Magos (central Portugal); an aspect of the octagonal chamber with open-top (OTC) (b); internal probes for CO₂ (left, down) and air temperature and humidity (right), a fan for air circulation (left, top), and a closed PVC chamber for measurement of greenhouse gases (GHGs) emissions inside the OTC (c).

an electric conductivity of 0.7 dS m⁻¹, low levels of mineral N, but a high level of chloride (71 mg Cl l⁻¹), and 30-48 mg Ca²⁺ l⁻¹, 51-87 mg Na⁺ l⁻¹, 7-10 mg K⁺ l⁻¹. ‘Ariete’ is a cultivar moderately sensitive to salinity and should not be negatively affected by the salts present in the water.

The experiments were kept free from weeds by using herbicides. The cultural practices used in the experiment were similar to the typical agricultural management used by Portuguese rice farmers for the last fourteen years.

Mean air temperature during rice growth in the open-field in 2012 varied from 18 °C in May to 21 °C in August (Figure 3). Minimum rainfall occurred in June-July (<10 mm) and maximum in October (100 mm). The wind speed ranged from 3.8 to 8.1 m s⁻¹ and global solar radiation presented a mean value of 5787 W m⁻². Mean, maximum

and minimum air temperatures inside the OTCs during the seasonal rice growth in 2012 are also shown in Figure 3. Mean air temperature variation ranged 2 °C in May and 3 °C in August. In summer, a higher number of days greater than 34 °C was registered compared with the open-field, and even several days above 38 °C (Figure 3). Inside the OTCs, solar radiation was reduced by less than 25% by the polyethylene film compared with the open-field.

Soil and floodwater sampling and analysis

Soil and floodwater samplings were taken in each plot, before and after fertilizer N application, at the end of flooding and at plant harvest (Figure 2). Soil samples were collected at 0-20, 20-40 and 40-60 cm depth, and analyzed for pH_(H2O) in a 1:2.5 soil: water suspension, mineral N determined colorimetri-

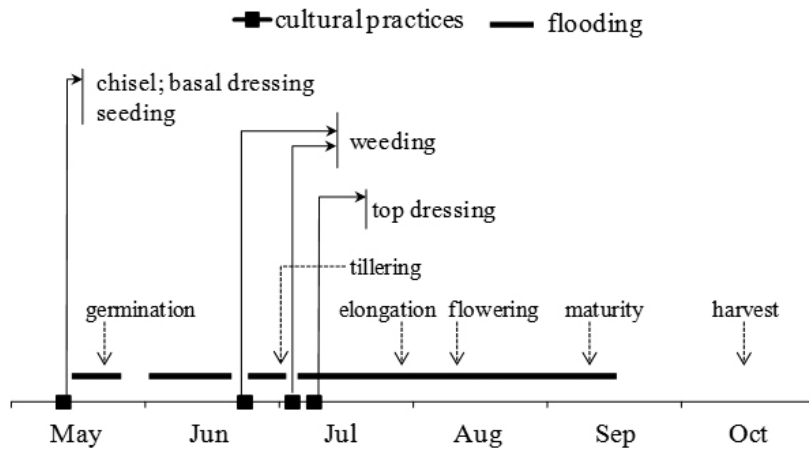


Figure 2 – Cultural practices and date of sampling in 2012, at Salvaterra de Magos (central Portugal).

cally (Skalar), and “fixed” NH_4^+ by adaptation of the method described by Silva and Bremner (1966). Briefly, 1.0 g soil ($<1.49 \mu\text{m}$) was treated with alkaline potassium hypobromite (KOB r -KOH) solution to remove exchangeable NH_4^+ and soil organic N. The clear supernatant liquid was decanted and discarded, while the residue was transferred into polyethylene centrifuge tubes and washed with 0.5 M KCl to remove residual exchangeable NH_4^+ . This operation was repeated three times. Finally, after decanting the clear supernatant liquid, the residue was treated with 5M HF-1M HCl solution and placed in an orbital shaker for 24 h to disintegrate the clay minerals containing the “fixed” NH_4^+ . The non-exchangeable form released by HF-HCl was determined by spectrometry. Floodwater sam-

ples were analyzed for pH and mineral N using the above reported methods.

Statistical analysis

Soil and floodwater results were analyzed by analysis of variance (ANOVA) through the General Linear Model (GLM) using the STATISTICA 6.0 software. Means separation was performed using the Bonferroni’s test for $p < 0.05$.

Results

Floodwater

In 2012, soil surface in the experimental field at Salvaterra de Magos (central Portugal) was in-

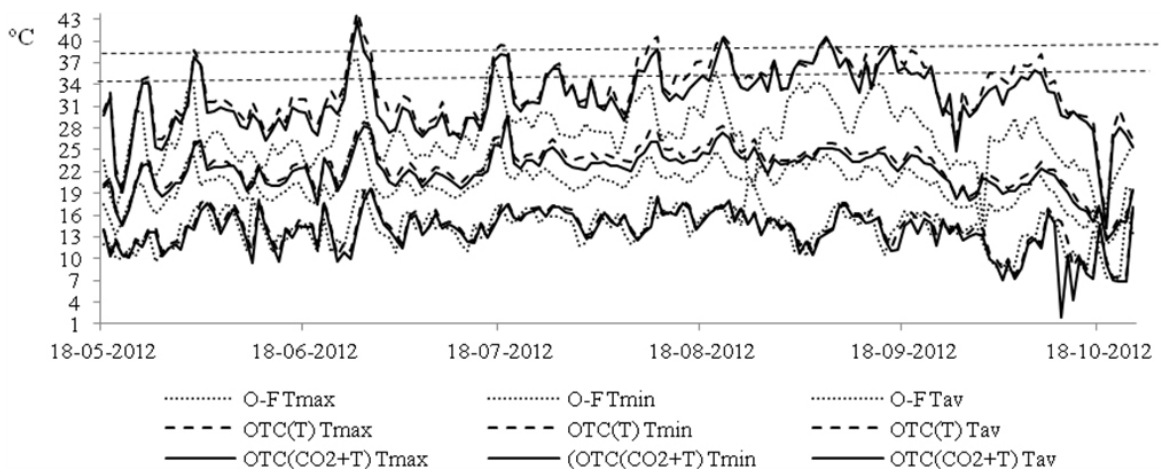


Figure 2 – Mean (Tav), maximum (Tmax) and minimum (Tmin) air temperatures in open-field (O-F) and inside the open-top chambers (OTCs) during the seasonal rice growth (2012), at Salvaterra de Magos (central Portugal).

termittently flooded after the crop sowing till the maturity stage (three weeks before the crop harvest). Treatments (open-field, temperature and CO₂+temperature elevation) did not affect the pH and mineral N in the flooded water, but the chemical properties varied significantly (p<0.001) during the growth season, especially after N fertilization (Table 1).

The pH in the floodwater increased during the season from pH=7.5 at initial flooding to pH=8.0 at maturity stage (Table 1). A higher NH₄⁺ level was measured after the topdressing, whereas a high NO₃⁻ content in floodwater was determined following the basal dressing+residue incorporation into soil.

Soil

In 2012, the dynamics of some chemical soil properties was evaluated during the seasonal rice growth (Table 2). The pH_(H₂O) and mineral N in the medium-textured clay loam soil varied significantly during the growth cycle, but were not affected by treatments (open-field, temperature and CO₂+temperature elevation). Non-exchangeable NH₄⁺ varied significantly with treatments, date of sampling and soil depth (Table 2).

Dynamics of soil pH

Soil pH_(H₂O) was slightly acid at the beginning of the experiment (pH_(H₂O)=6.1) and decreased signifi-

Table 1 – ANOVA (MLG) results and means separation for pH and soluble NH₄⁺-N and NO₃⁻-N in floodwater during the seasonal growth of rice ‘Ariete’ in 2012, at Salvaterra de Magos (central Portugal).

	pH	NH ₄ ⁺ -N (mg l ⁻¹)	NO ₃ ⁻ -N (mg l ⁻¹)
Treatment			
Open-field	7.5a	0.34a	0.52a
CO ₂ +temperature	7.6a	0.39a	0.50a
Temperature	7.5a	0.31a	0.53a
Date of sampling			
22 May	7.5bc	0.02b	0.52b
30 May	6.7d	0.21b	1.66a
4 July	7.3c	0.18b	0.14e
14 July	7.7b	1.46a	0.16de
8 August	8.0a	0.02b	0.33c
12 September	8.0a	0.21b	0.31cd
Source of variation	F-value	F-value	F-value
Treatment (T)	ns	ns	ns
Date of sampling (D)	66.61***	51.5***	218.1***
Interactions:			
T x D	6.80***	ns	3.65**

ns, **, ***= F-values, non-significant (p>0.05) and significant at p<0.01 and p<0.001, respectively according to Bonferroni's test.

Table 2 – ANOVA (GLM) results and means separation for pH(H₂O), available mineral-N (NH₄⁺, NO₃⁻) and non-exchangeable NH₄⁺ during the seasonal growth of rice ‘Ariete’ in 2012, at Salvaterra de Magos (central Portugal) in response to treatments.

	pH _(H₂O)	NH ₄ ⁺ -N (mg kg ⁻¹)	NO ₃ ⁻ -N (mg kg ⁻¹)	Non-exchangeable NH ₄ ⁺ -N (mg kg ⁻¹)
Treatment				
Open-field (F)	5.9a	4.6a	2.38a	258.1a
CO ₂ + temperature (C)	5.9a	4.8a	2.54a	224.7b
Temperature (T)	5.9a	4.8a	2.21a	217.9b
Date of sampling				
14 May	6.1a	6.6b	2.59bcd	182.5de
16 May	5.6c	8.7a	5.37a	321.2a
30 May	5.8c	3.5cd	3.54b	227.1cd
26 June	5.8c	5.1bc	2.55bcd	256.5bc
13 July	5.8c	6.7ab	1.35cde	287.9ab
8 August	6.1a	3.5cd	0.18e	171.7e
12 September	6.0ab	0.7e	0.80de	225.6cd
10 October	6.0ab	3.1d	2.63bc	196.0de
Soil depth (cm)				
0-20	5.9a	4.9a	4.59a	247.5a
20-40	5.9a	4.5a	1.56b	234.8ab
40-60	5.9a	4.8a	0.99b	218.4b
ANOVA				
	F-value	F-value	F-value	F-value
Treatment (T)	ns	ns	ns	11.02***
Date of sampling (D)	12.8***	33.3***	16.5***	24.28***
Soil depth (d)	ns	ns	61.3***	5.02**
T x D	1.9*	ns	ns	2.74**
D x d	ns	2.6**	14.5***	ns

ns, *, **, ***= F-values, non-significant (p>0.05) and significant at p<0.05, p<0.01 and p<0.001, respectively according to Bonferroni's test.

cantly during the season till the topdressing with sulfamid application ($\text{pH}_{(\text{H}_2\text{O})}=5.8$). After the topdressing, soil pH increased again to about $\text{pH}=6.0$, and was maintained constant until the crop harvest (drained soil). Soil pH did not vary with soil depth (Table 2).

Dynamics of soil mineral N

Available N in soil increased significantly after the basal dressing and residue incorporation into the soil to 8.7 and 5.4 mg kg^{-1} , respectively for NH_4^+ and NO_3^- for the mean effect of treatment and soil depth (Table 2), but decreased thereafter to the crop maturity (flooded soil) and harvest (drained soil), with insignificant levels at these stages. While NH_4^+ did not vary with soil depth (4.8 mg kg^{-1}), NO_3^- was significantly reduced in the profile from about 5 to 1 mg kg^{-1} .

Dynamics of non-exchangeable NH_4^+

A higher concentration of non-exchangeable NH_4^+ ($p<0.001$) was measured in the open-field (258 $\text{mg NH}_4^+\text{-N kg}^{-1}$) for the mean effect of soil depth, compared with the OTCs, which did not vary significantly from each other (221 $\text{mg NH}_4^+\text{-N kg}^{-1}$). A “native fixed” NH_4^+ amounting to 183 mg kg^{-1} was determined in the bulk soil prior to basal dressing and flooding, 28 times greater than the available NH_4^+ in the soil at that time. Thereafter, after the basal dressing+residue incorporation+flooding, a significant rise on “fixed” NH_4^+ was verified (321

$\text{mg NH}_4^+\text{-N kg}^{-1}$) and again after the topdressing to a similar amount (288 $\text{mg NH}_4^+\text{-N kg}^{-1}$). The level of “fixed” NH_4^+ in soil after the basal dressing corresponded to about 37 times the concentration of NH_4^+ in the soil solution at this time, but the proportion was 43 times greater after the topdressing. Following the topdressing, the concentration of “fixed” cation decreased significantly thereafter to the same level as that at the start of the experiment in the OTCs (Figure 4), but was higher ($p<0.01$) in the open-field (280 $\text{mg NH}_4^+\text{-N kg}^{-1}$).

Non-exchangeable NH_4^+ was significantly lower in the 40–60 cm layer, decreasing from a mean value of 241 $\text{mg NH}_4^+\text{-N kg}^{-1}$ at 0-40 cm layer to 218 $\text{mg NH}_4^+\text{-N kg}^{-1}$ (Table 2). In the present clay loam soil, clay content did not vary significantly with soil depth (534-564 g kg^{-1} soil). Therefore, the higher “newly fixed” NH_4^+ in the 0-40 cm depth can be attributed to the organic residue+mineral N fertilizer incorporated at surface soil.

Discussion

Rice residues (straw+root) were partially burnt in November 2011 in the field and incorporated into the clay loam soil simultaneously with the basal mineral N, before rice sowing in 23 May 2012, followed by flooding. The incorporation of residues and flooding favored deamination of soil organic N by the facultative anaerobic microorganisms in

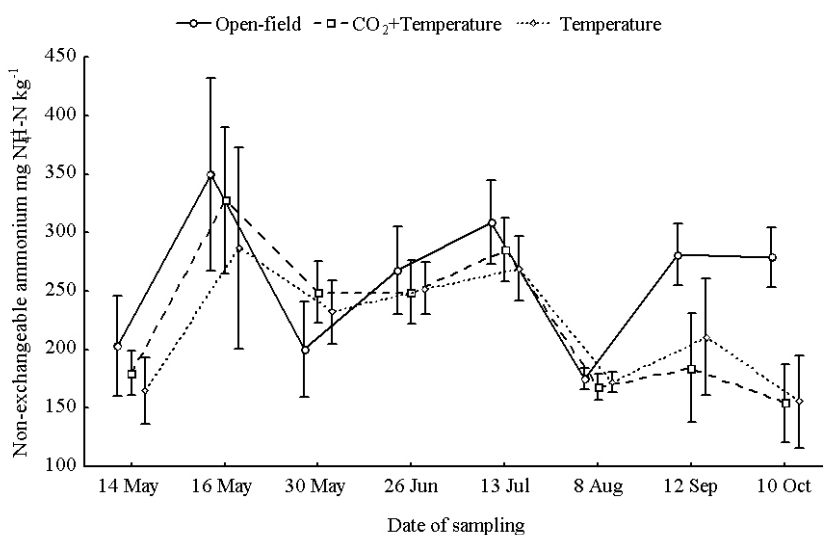


Figure 4 – Mean (T_{av}), maximum (T_{max}) and minimum (T_{min}) air temperatures in open-field (O-F) and inside the open-top chambers (OTCs) during the seasonal rice growth (2012), at Salvaterra de Magos (central Portugal).

the anaerobic conditions. Ammonium was then the predominant N form in these conditions remaining in the soil at higher levels than NO_3^- (Table 2). Nitrification was higher in the surface layer, in the rhizosphere (Table 2). The frequent application of 5-20 cm flood irrigation produced an almost constant anaerobic condition in lower layers, except the surface layer due to oxygenation by the water itself, the wind movement, and the roots aerenchyma (Zhang and Scherer, 2002).

Nevertheless, in our study overall available NH_4^+ in the soil was 28-43 times smaller than the overall "fixed" cation which accounted for an average of 11% of total N in the plough layer, in agreement with the literature. Just two days after flooding and application of basal N dressing, NH_4^+ "fixation" rose by 76% (Figure 4). This fact was repeated sixteen days after the N topdressing under flooding conditions with a 58% increase in relation to the bulk soil. When subtracting "native fixed" NH_4^+ from apparent "fixed" NH_4^+ after N fertilization, the net increase in "fixed" NH_4^+ was larger after the basal dressing in drained soil (+146 mg $\text{NH}_4^+\text{-N kg}^{-1}$, corresponding to an average 322 kg N ha^{-1} in the 0-20 cm layer) than flooded soil after topdressing (+105 mg $\text{NH}_4^+\text{-N kg}^{-1}$ corresponding to an average 232 kg N ha^{-1} in the same layer). These rates were greater than the amount of mineral N added to the soil in each growth stage (60 kg N ha^{-1}) and it is assumed to be a result of the addition of mineral N, the higher NH_4^+ concentration in soil solution by mineralization of organic N compounds, and by flooding with a probable reduction of Eh (not determined) followed by an increase in the surface layer of clay minerals. Soil $\text{pH}_{(\text{H}_2\text{O})}$ decreased significantly (0.3 pH units) after initial flooding, indicating a Eh reduction. Especially during the first week after flooding, Zhang and Scherer (2002) also observed that the concentration of non-exchangeable NH_4^+ was significantly higher. Kowalenko and Cameron (1976) verified that approximately one half of the added NH_4^+ in the form of $(\text{NH}_4)_2\text{SO}_4$ was fixed within two days after addition to the incubated clay loam waterlogged soil.

Under waterlogging and two weeks after the basal dressing or four weeks after the topdressing, the concentration of "fixed" NH_4^+ declined significantly in the our study. Zhang and Scherer (2002) observed a similar situation two weeks after fertilization, but in the contrast to the present situation where the non-exchangeable NH_4^+ was higher in the 0-40 cm layer, these same authors found that on the rhizosphere the "fixed" cation was reduced because it was released in the rhizosphere where the Eh

was increased by the oxygen (O_2) emitted by roots (but this diffusion decreased with distance from roots). According to Barshad (1954) and Newman (1969) the release of non-exchangeable NH_4^+ is favored by protons. Protons may penetrate into the wedge zones due to the somewhat expanded state of the clay minerals and displace the "specifically adsorbed" cations (Sparks and Liebhardt, 1982). In the present study, the reason for diffusion of "fixed" cation into the solution in 30 May and 8 August 2012 was not clear but apparently was related to an expanding of clay minerals and a depletion of available NH_4^+ in the soil solution (Table 2).

Non-exchangeable NH_4^+ was significantly higher in the open-field conditions compared with the climate change treatments in OTCs (Table 2). The interaction of treatments with seasonal date of sampling revealed that this difference was relevant from 8 August 2012 until harvest. Treatments for climate change in OTCs did not affect the dynamic of non-exchangeable NH_4^+ and other chemical properties in the soil. Maximum daily temperature in summer, from middle July to middle September 2012 was above 34 °C in the OTCs (Figure 3), whereas in the open-field maximum temperature was below 34 °C. This fact can explain the lower NH_4^+ "fixation" under the maximum daily temperature greater than 34 °C inside the OTCs which agrees with some literature that refers that "fixation" is stronger under 30 °C (Jia *et al.*, 2000).

Few references on the effects of CO_2 elevation on soil N dynamics are available (French *et al.*, 2009), but on the NH_4^+ "fixation" they were not found. In the present study, the CO_2 elevation simultaneous with the temperature did not affect the dynamics of non-exchangeable cation. A similar response was found for soil microbial activity (data not shown). In addition, the increase of CO_2 concentration promoted photosynthesis, plant growth, and root exudates but together with the temperature elevation the crop yield was significantly reduced compared with the production under the open-field (data not shown). Further long-term field studies are recommended for the effect of global climate change on "fixation" of NH_4^+ , microbial communities and dynamics of C and N in soil.

Conclusions

Non-exchangeable NH_4^+ was significantly higher under open-field conditions compared with the evaluated climate change scenarios of elevation of temperature with or without elevated atmospheric

[CO₂]. Temperature elevation reduced the “fixation” rate of the cation, while the CO₂ concentration rise did not affect particularly the non-exchangeable form. Further studies are required for consolidation of these findings together with microbial communities and dynamics of C and N in soil.

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